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 [EE] Evening Poster | A (Atmospheric and Hydrospheric Sciences) | A-GE Geological & Soil Environment

## [A-GE31]Subsurface Mass Transport, Material Cycle, and Environmental Assessment

convener:Yuki Kojima(Department of Civil Engineering, Gifu University), Shoichiro Hamamoto(Department of Biological and Environmental Engineering, The University of Tokyo), Hiroataka Saito(東京農工大学大学院農学研究院, 共同), Yasushi Mori(Graduate School of Environmental and Life Science, Okayama University)

Mon. May 21, 2018 5:15 PM - 6:30 PM Poster Hall (International Exhibition Hall7, Makuhari Messe)

This session covers the topics on mass transport, water and energy cycles in geoenvironment. Subjects related to laboratory and field measurements, theoretical analysis, and numerical modeling will be discussed. Presentations on geo-pollution, remediation, geological disposal of hazardous wastes, ground source heat utilization, mass transport in vadose zone, soil-water monitoring, and environmental assessment are encouraged.

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## [AGE31-P05]Evaluation of $^{137}\text{Cs}$ bioavailability decay in French pastures from long-term field monitoring

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Keywords:137Cs, soil, plants, bioavailability, modelling

Soil-to-plant transfer of  $^{137}\text{Cs}$  to grass and crops is a crucial pathway since this is considered as a long-lasting input of radioactivity into the human food chain. A long-term radiological monitoring of remaining artificial radionuclides from atmospheric nuclear tests and Chernobyl accident has been conducted by IRSN since the 1990's at a dozen of pasture plots located in France. The sampling sites are located in areas heavily affected by the atmospheric depositions, but also in the environment of nuclear facilities. The monitoring involves measurements of  $^{137}\text{Cs}$  activity in upper soil (0-5 cm), grass and cattle milk, continuously sampled at the same plot, leading to a total of about 600 data available for analysis. The observed values of the soil-to-plant transfer factor ('field Tf'), defined as the ratio between  $^{137}\text{Cs}$  activities in vegetation and soil, vary by more than one order of magnitude from one site to another, with a mean value which is significantly lower than those given by IAEA (2010) (0.29 and 0.18 for sandy and clayed soils, respectively). It is often assumed that environmental parameters such as soil physical and chemical properties (clay content, but also pH, organic matter, exchangeable potassium, etc.) or plants species can explain a part of the variability observed between sites. The issue of the present research is to compare observed Tfs with theoretical values predicted from semi-mechanistic models derived mainly from experiments and available in the literature (Smolders et al., 1997; Absalom et al., 1999; Absalom et al., 2001; Tarsitano et al., 2011). 'Predicted Tf' requires the determination of empirical parameters such as the labile distribution coefficient K<sub>dl</sub> at the soil-soil solution interface and the concentration factor CF, at the soil solution-roots interface. The calculation of K<sub>dl</sub> and CF requires knowledge of the soil physico-chemical parameters cited before. An additional dynamic parameter (namely Dt) is needed which represents the ageing of  $^{137}\text{Cs}$  in the rooting layer, defined as the percentage of bioavailable  $^{137}\text{Cs}$  with respect to time. It is usually assumed that Dt obeys a kinetics equation involving two kinetic rates, i.e. a fast and a slow components denoted  $k_{\text{fast}}$  ( $1.9 \cdot 10^{-3} \text{ d}^{-1}$ ) and  $k_{\text{slow}}$  ( $1.9 \cdot 10^{-4} \text{ d}^{-1}$ ) respectively, which values must be ascertained (Absalom et al., 1999). Special attention should be paid to the calculation of Dt and hence on the chosen values of  $k_{\text{fast}}$  and  $k_{\text{slow}}$  which drastically influence the 'predicted Tf' on the long term. Thus, before any comparison between 'predicted Tf' and 'field Tf', an attempt is made in the

present study to evaluate  $k_{\text{slow}}$  using our time observations of  $^{137}\text{Cs}$  activities in grassland plants and milk.  $k_{\text{fast}}$  could not be evaluated because the monitoring started in the 90's, several years after deposition. The long-term time series of  $^{137}\text{Cs}$  activity in vegetation and milk recorded at 13 plots of the French territory allows us to deduce an apparent decay rate for each plot,  $I_{\text{app}}$  ( $\text{d}^{-1}$ ), from which the corresponding decay rate of  $^{137}\text{Cs}$  activity in the upper soil layer,  $I_{\text{mig}}$  ( $\text{d}^{-1}$ ), was subtracted. This enables to estimate the slow component of the bioavailability factor  $Dt$ , i.e.  $k_{\text{slow}} = I_{\text{app}} - I_{\text{mig}}$ , with two distinct values at each site (for grass and milk). The observed median values of  $k_{\text{slow}}$  are  $1.6 \cdot 10^{-4}$  and  $1.7 \cdot 10^{-4} \text{ d}^{-1}$  for respectively grass and milk. These are close to values deduced from data acquired in Switzerland (four monitoring plots) using the same methodology, namely  $1.3 \cdot 10^{-4}$  and  $2.3 \cdot 10^{-5} \text{ d}^{-1}$  (Corcho-Alvarado et al., 2016). Our results thus corroborate the time constant used in the  $Tf$ 's equations according to Absalom et al. (1999). Thus in the following of this research project a site specific value of  $Dt$  will allow us to estimate  $Tf$  for each monitored plot.