[JJ] Evening Poster | M (Multidisciplinary and Interdisciplinary) | M-TT Technology & Techniques

## [M-TT37]Frontiers in Geochemistry

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Sun. May 20, 2018 5:15 PM - 6:30 PM Poster Hall (International Exhibition Hall7, Makuhari Messe) Many new findings in earth and planetary sciences have been obtained by using state-of-the-art techniques supported by new technical development in analytical chemistry. This session aims at providing an opportunity for those developing new analytical methods to get together and have a strategic discussion on frontiers in geochemistry and cosmochemistry. We welcome a wide range of cutting-edge geochemical topics based on technical development, which have a potential for breakthrough of earth and planetary sciences. Besides, topics related to the direction of geochemistry and cosmochemistry in future are also welcome. Especially, we welcome topics which present how to install/maintain precious facilities in geochemical laboratories. We expect wide-ranged and futureoriented discussion to develop geochemistry and cosmochemistry.

## [MTT37-P01]Determination on the triple oxygen isotopic compositions of tropospheric ozone

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Tropospheric ozone  $(O_3)$  is important as a greenhouse gas, as well as having harmful effects on respiration and photosynthesis. In addition,  $O_3$  is important as an oxidant in the tropospheric photochemical reactions. In recent years, tropospheric ozone have been increasing in Eastern Asia, and thus we must understand the origin and behavior of tropospheric ozone accurately.

In this study, we determined the oxygen isotopic compositions including the triple oxygen isotopic compositions (Δ<sup>17</sup>O), by passing air sample through nitrite (NO<sub>2</sub><sup>-</sup>)-coated filters, which allows the reaction of O<sub>3</sub> with NO<sub>2</sub><sup>-</sup> to produce NO<sub>3</sub><sup>-</sup>. The oxygen isotopic composition of NO<sub>3</sub><sup>-</sup> was then determined to estimate the oxygen isotopic composition of O<sub>3</sub>. To remove atmospheric HNO<sub>3</sub> prior to the reaction between O<sub>3</sub> and NO<sub>2</sub><sup>-</sup>, nylon filters were placed before the nitrite-coated filters. In order to obtain a high collection efficiency for the reaction between O<sub>3</sub> and NO<sub>2</sub><sup>-</sup>, the O<sub>3</sub> collection was conducted at a flow rate less than 0.5L/min. The NO<sub>2</sub><sup>-</sup> on the filter was removed by adding hydrogen azide (N<sub>3</sub>H) which selectively converts NO<sub>2</sub><sup>-</sup> to N<sub>2</sub>O. Then, NO<sub>3</sub><sup>-</sup> was converted to N<sub>2</sub>O, which is converted to O<sub>2</sub> to be injected into an isotope mass spectrometer (MAT252), allowing quantification on the &Delta;<sup>17</sup>O value of NO<sub>3</sub><sup>-</sup>. From the values determined for NO<sub>3</sub><sup>-</sup>, the oxygen atoms derived from NO<sub>2</sub><sup>-</sup> was subtracted to determine the oxygen isotopic compositions of O<sub>3</sub>. Please note that the oxygen atoms in O<sub>3</sub> (&Delta;<sup>17</sup>O (O<sub>3</sub>) <sub>bulk</sub>), but is the isotope composition of the oxygen atoms in the terminal positions (&Delta;<sup>17</sup>O (O<sub>3</sub>) <sub>terminal</sub>) of O<sub>3</sub>.

Observation on the atmospheric O<sub>3</sub> was conducted from August to December, 2017, at Nagoya University. The &Delta;<sup>17</sup>O values of ozone were between +32&permil;~+39&permil; which coincided well with those previously determined for tropospheric ozone (35&permil;&plusmn;4&permil;). Besides,

the Δ<sup>17</sup>O values of ozone were the lowest in August, and were the highest in November. The seasonal variation in the &Delta;<sup>17</sup>O values is most likely due to the stratospheric influence on the tropospheric  $O_3$ . We also found about 1‰ differences in the Δ<sup>17</sup>O values between day and night. We concluded that the formation of an inversion layer in night time was responsible for the lower &Delta;<sup>17</sup>O values. That is to say, while the <sup>17</sup>O-depleted  $O_3$  produced at ground level heights under the high pressure condition occupied major portion of  $O_3$  in night time due to the inversion layer, the <sup>17</sup>O-enriched  $O_3$  produced at the upper layers contributed to  $O_3$  in day time through the active vertical convection.