## Raman spectroscopic investigation of $\alpha$ - $\beta$ quartz phase transition in hydrothermal diamond-anvil cell and acquisition of equation of states of aqueous solutions

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The  $\alpha$ - $\beta$  quartz phase transition temperatures ( $T_{tr, qtz}$ ' s; up to ~781 °C) were determined at various pressures (P' s; up to ~802 MPa) in a hydrothermal diamond-anvil cell (HDAC; Li et al., 2016; Rev. Sci. Instrum., **87**, p. 053108-1) by monitoring the Raman shift of the  $\alpha$ -quartz band near 128 cm<sup>-1</sup> (at 24.2 °C). When compared with the commonly used  $\alpha$ -quartz band near 465 cm<sup>-1</sup> (e.g., Schmidt and Ziemann, 2000; Am. Mineral. **85**, p. 1725), the rate of reduction in wavenumber during H<sub>2</sub>O isochoric heating is about five times higher. In addition, the sudden change of the rate of reduction in wavenumber at  $T_{tr, qtz}$  is much more sharp and clear, making the  $\alpha$ -quartz band near 128 cm<sup>-1</sup> a much better choice for detecting the phase transition.

Our experimental procedures were similar to those of Shen et al. (1993; Am. Mineral. **78**, p. 694), except Raman spectroscopic method instead of laser interferometry was used to determine  $T_{tr, qtz}$ . A quartz wafer (prepared from a natural crystal from Asikaerte Be pegmatite in Xinjiang, China) together with or without  $H_2O$  were loaded in the sample chamber, which was a hole in a Re gasket between two diamond anvils;  $H_2O$  pressure medium was not needed for experiments at 0.1 MPa total pressure.

Raman spectra were acquired during a heating cycle of the experiments after the sample was kept at a constant *T* for more than 3 minutes. We used a JY/Horiba LabRam HR Evolution Raman system, with 532.06 nm (frequency doubled Nd:YAG) laser excitation, a SLWD 50x Olympus objective having 0.35 numerical aperture, a 1800-groove/mm grating with a spectral resolution of about 0.2 cm<sup>-1</sup>, and ~14 mW laser light was focused on the sample during the measurement. Spectra were collected in one spectrographic window (from 77 to 593 cm<sup>-1</sup>) for either 30 s (below 700 °C) or 60 s (above 700 °C) with two accumulations per spectrum.

The bulk density of H<sub>2</sub>O in the sample chamber for the observed  $T_{tr, qtz}$  was determined by measuring the homogenization  $T(T_h)$  after the liquid-vapor phase separation during isochoric cooling. The two K-type thermocouples in HDAC were calibrated with the melting points of NaNO<sub>3</sub> (306.8 °C) and NaCl (800.5 °C), and the uncertainties in *T* measurements are ±1.5 °C. The associated pressures at  $T_h$  ( $P_h$ ) and  $T_{tr, qtz}$  ( $P_{tr, qtz}$ ) were calculated based on the equation of state (EOS) of H<sub>2</sub>O (Wagner and Pru $\beta$ , 2002; J. Phys. Chem. Ref. Data **31**, p. 387). The straight line connecting ( $T_h$ ,  $P_h$ ) and ( $T_{tr, qtz}$ ,  $P_{tr, qtz}$ ) in a *P*-*T* space is near the isochore of that bulk density of H<sub>2</sub>O. Similar approach was successfully applied to obtain isochores of 2 m ZnCl<sub>2</sub> solution (Bassett et al., 2000; Zeitsch. Kristallogr. **215**, p. 711), and will be extended to other geologically important aqueous solutions at *T* 's up to 1000 °C using HDAC and Raman spectroscopy. The  $\alpha$ - $\beta$  quartz phase boundary obtained in this study can be represented by: ( $P_{tr, qtz}$ ) (±8.8 MPa) = 0.0015 ( $T_{tr, qtz}$ )<sup>2</sup> + 1.8268 ( $T_{tr, qtz}$ ) –1544.5, where ( $T_{tr, qtz}$ ) is between 574 and 781 °C; with R<sup>2</sup> = 0.9998. Our results agree, within experimental uncertainties, with those reported by Mirwald and Massonne (1980; J. Geophys. Res. **85**, p. 6983), but with some deviations from other previous data.

Keywords:  $\alpha - \beta$  quartz phase transition, Raman spectroscopy, Equation of states of aqueous solutions, Hydrothermal diamond-anvil cell, Isochore SIT26-12

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